Seasonal	Diurnal	Variation	of	Η,	<i>1922</i> .
(Coeffic	cients of	the first fo	ur	teri	ns.)

	P ₁ .	A ₁ .		P ₂ .	A ₂ .		P ₃ .	A ₃ .		P ₄ .	A4.	
Four equinoctial months	 γ 13.43 9.78 7.30	° 291 257 208	37 00 55	γ 8·86 4·94 3·01	299 275 224	, 32 19 19	γ 4·52 3·14 1·97	303 285 202	, 43 34 42	γ 0·35 0·59 1·17	276 248 186	55 43 19

The summer months taken are January, February, November, December; equinoctial months taken are March, April, September, October; winter months taken are May, June, July, August. As regards the amplitudes of these harmonics, a general seasonal graduation shows itself in the first three terms, the amplitude being greatest in summer-time; but in P_4 the reverse is the case, the amplitude being greatest in winter and very small in summer; but there seems no reason to doubt that it is well determined, so that some significance attaches to its phase angle differing by very approximately 90° between summer and winter. Significant also is the fact that in the third and fourth components the winter phase angle is $\frac{2}{3}$ of the summer angle phase, very nearly. In A_2 the summer angle is practically $\frac{4}{3}$ of the winter value, so that in these harmonics the summer angle is $\frac{4}{3}$ of the winter angle for A_4 and A_3 .

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There seems to be more evidence than usual of the non-independence of the harmonic constituents. This may be looked upon as worthy of consideration, and perhaps attributable to our latitude. The origin of the averages must be remembered, for the summer data are the means for two intervals of two months each, separated by ten months; for the equinoctial data the separation is six months; for the winter data, really two months; and the mean values from the three means is this year appreciably different, being—

II, summer months.. 0.222209 C.G.S.

H, equinoctial 0·222125 H, winter 0·222162

Year 0.222165

The winter value is very near the arithmetical mean for the year.

Coefficients of the First Four Terms in Fourier Series, of the Main Seasonal Diurnal Variation, for Year 1922 and for Greenwich Civil Mean Day, at Christchurch, N.Z.

		P ₁ .	A ₁ .		P ₂ .	A	2.	P ₃ .	A	}•	P ₄ .	Α,	1•
	M	l agne	tic H	rizon	tal For	ce (u	nit 1	γ).					
			•	,		0	,		0	,		0	,
Four summer months	113	3.43	291	37	8.86	299	32	4.52	303	43	0.35	262	57
Four equinoctial months		9.78	257	00	4.94	275	19	3.14	285	34	0.55	272	14
Four winter months		7.30	208	55	3.01	224	19	1.97	202	42	1.17	186	19
Four winter months													
Year	8	8.61	261	51	5.02	282	23	5.84	334	06	0.57	226	16
		Maa	metic	Decl	ination	· (111	rit 1')					
						•	•		. 00	90	0.10	1 40	4 ~
Four summer months		2.59	14	09	2.19	357	48	0.20	22	$\frac{39}{50}$	0.16	43	45
Four equinoctial months	I	1.28	35	23	1.83	339	46	0.79	347	58	0.25	332	12
Four winter months	[-	5.06	170	54	0.97	328	5 4	0.43	338	16	0.27	289	15
Year		1.36	32	01	1.63	345	33	0.84	00	33	0.17	329	06

The phase angles are as at 0h. G.C.M.T.

The four summer months comprise January, February, November, December; the equinoctial months, March, April, September, October; and the four winter months, May, June, July, August. In the horizontal force it is noticeable that the amplitude of the diurnal wave P_1 in summer seems to determine all the amplitudes during the equinoctial months; and in the same way the amplitude of the diurnal wave in equinox seems to determine all the amplitudes in the winter months. This is possibly dependent upon the consideration that the amplitude of P_1 depends upon a generally effective cause. Such cause may be expected to have an effect upon the amplitude of the constituent waves P_1 , P_2 , P_3 , P_4 in the ratio 1, $\frac{1}{2}$, $\frac{1}{3}$, $\frac{1}{4}$. And taking the resolved equinoctial amplitudes in H, we find $9.78 + \frac{1}{2}(4.94) + \frac{1}{3}(3.14) + \frac{1}{4}(0.55) = 9.78 + 2.47 + 1.05 + 0.14 = 13.44r$, which is the value of P_1 in summer. Similarly, for the winter months, in H, we find $7.30 + \frac{1}{2}(3.01) + \frac{1}{3}(1.97)$